

# Wind regimes above and below a dense oil palm canopy: Detection of decoupling and its implications on CO<sub>2</sub> flux estimates

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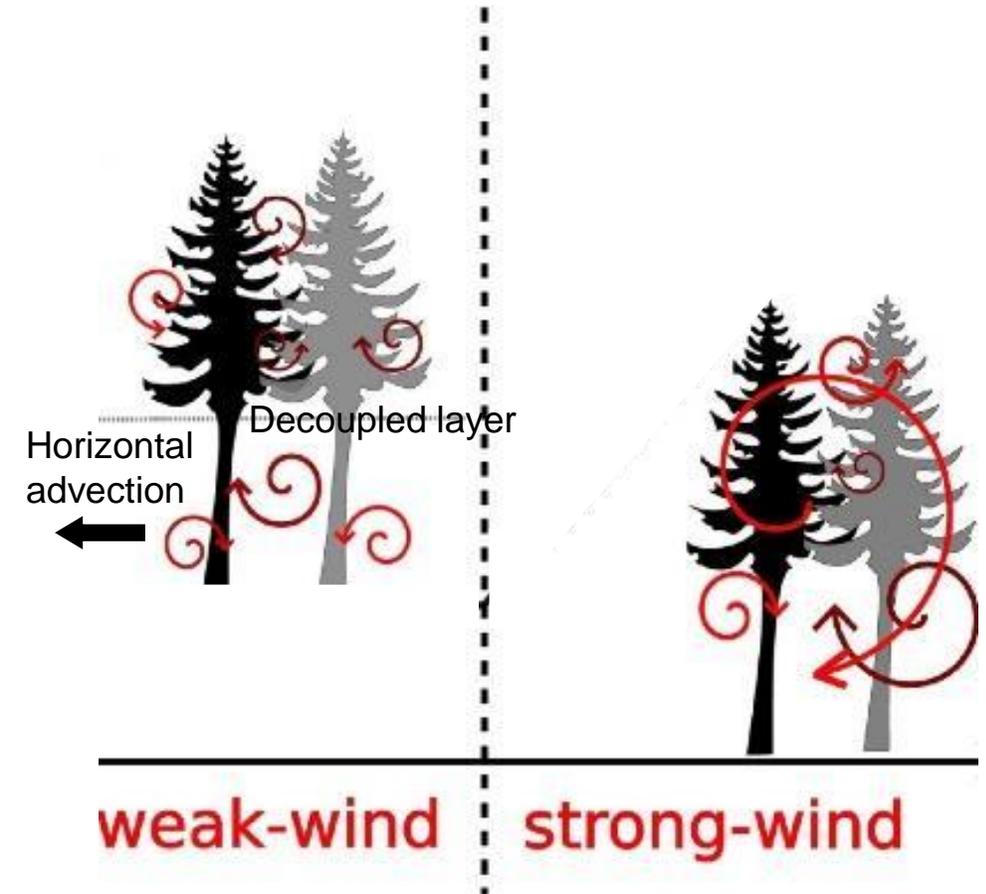
**DFG** Deutsche  
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# Decoupling

- Especially during night, calm weather conditions in tall vegetation canopies, such as forest or oil palm, may result in the formation of an isolated layer near the surface, which is decoupled from the above-canopy air layer.
- When decoupling occurs, there is a high potential that above-canopy measured carbon dioxide ( $\text{CO}_2$ ) based on eddy covariance (*EC*) measurements might not represent the true ecosystem  $\text{CO}_2$  flux as below-canopy respiration might be undetected by the *EC* system.

## Schematic representation of air mixing within tall vegetation canopies



Adapted from Freundorfer et al., 2019,  
*Agricultural and Forest Meteorology* 279.

# Study aim

- **Investigate wind dynamics** of a mature oil palm (*Elaeis guineensis* Jacq.) plantation in tropical lowland Jambi Province (Sumatra, Indonesia).
- **Assess the strength of turbulent mixing** as an estimator for the degree of above- and below-canopy coupling by using eddy covariance (*EC*) measurements.
- **Explore** the potential **implications of decoupling** and horizontal below-canopy flow on the above-canopy derived net ecosystem exchange (*NEE*).
- Explore the **characteristics of vertical CO<sub>2</sub> concentration** in the oil palm plantation.

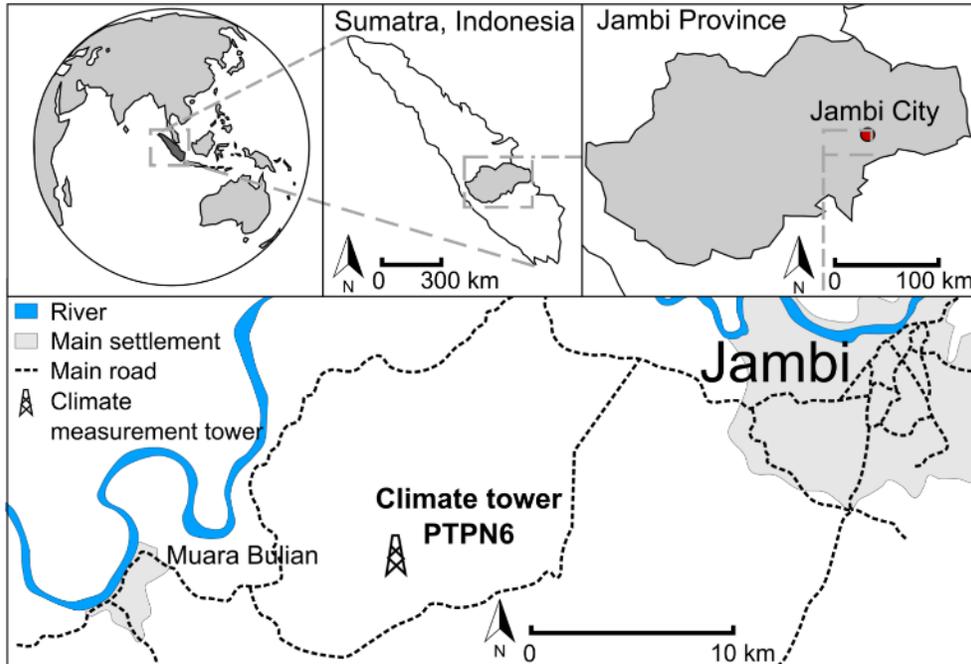


EC-tower within the studied oil palm plantation.

# Study site

Our study has been conducted in a mature commercial oil palm plantation in tropical lowland Jambi Province (Sumatra, Indonesia). Palms were planted in 2002, with 156 palms per hectare, and reach an average height of 12 meters.

## Study location



Video of oil palm plantation and measurement tower

Stiegler et al., 2019, *Biogeosciences* 16.

Drescher et al., 2016, *Phil. Trans. R. Soc. B* 371.

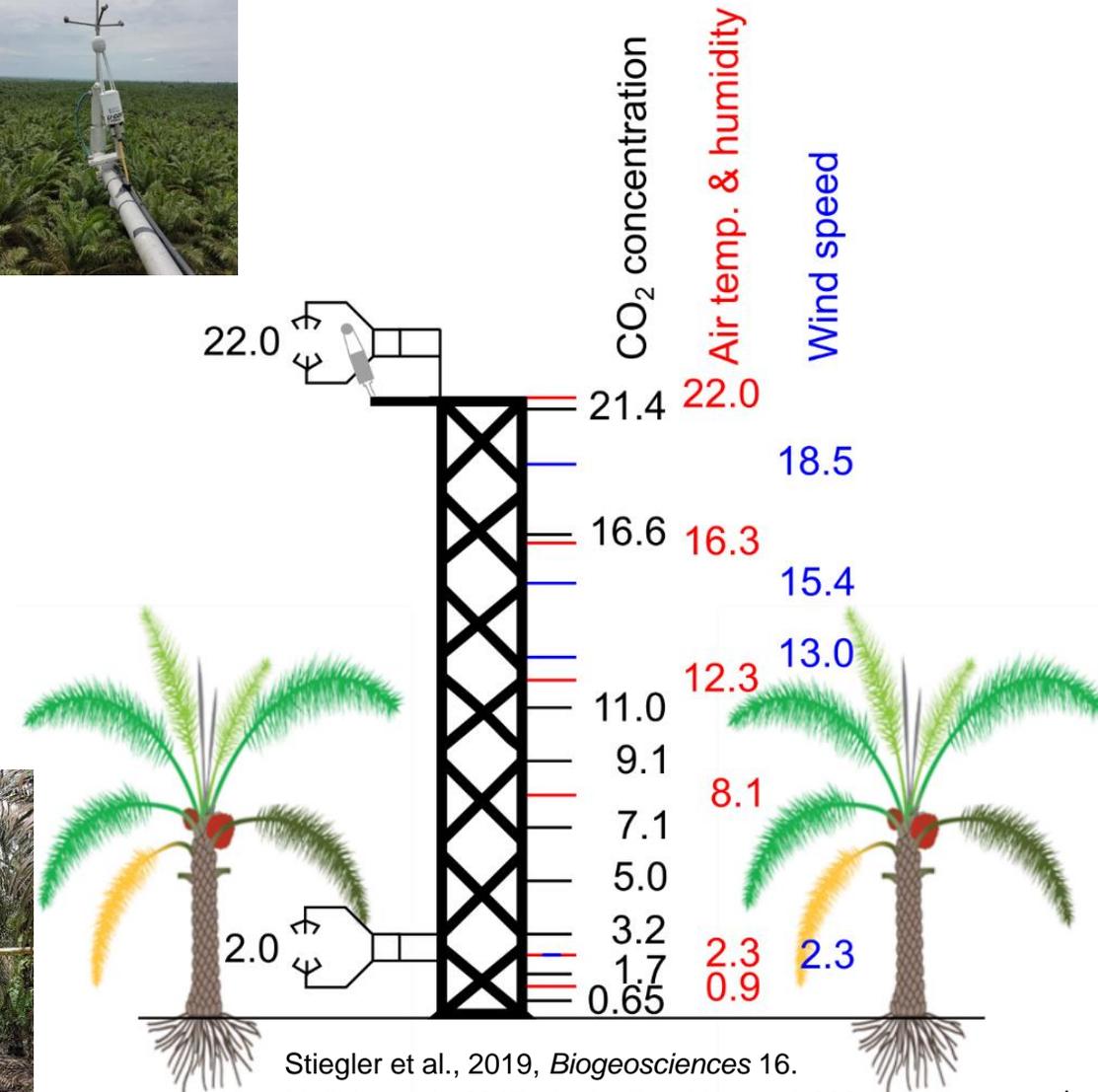
Clough et al., 2016, *Nature Communications* 7.

# Data collection & processing; Identification of decoupling & flux filtering

- Flux data were obtained from August 2017 to July 2020 by two eddy covariance (*EC*) systems
- Three approaches based on *EC* derived wind measurements to identify decoupling:
  1. Threshold based on friction velocity ( $u^*$ )
  2. Threshold based on the correlation between the standard deviation of below vs. above-canopy vertical wind speed ( $\sigma_w$ )
  3. Threshold based on dynamic stability of the atmosphere (bulk Richardson number,  $Ri_b$ )
- Thresholds were derived with R-package "*strucchange*". Filtered *EC-data* was gapfilled using REddyProc (MPI Jena, Germany) to derive cumulative CO<sub>2</sub> fluxes
- Meteorological parameters and CO<sub>2</sub> concentration were measured at various heights along the tower



Climate tower & measurement heights



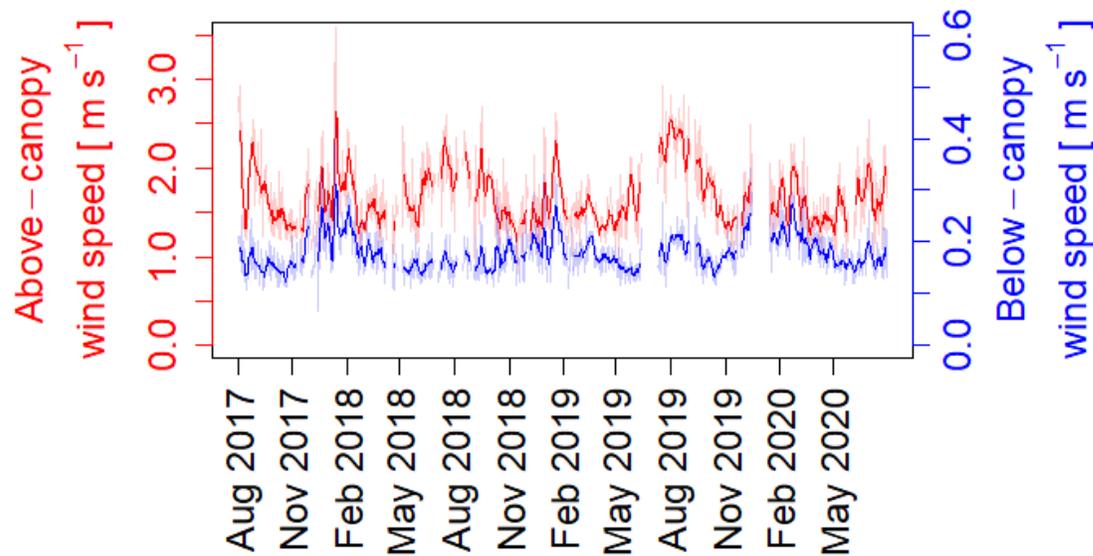
Stiegler et al., 2019, *Biogeosciences* 16.

Meijide et al., 2017, *Agric. For. Meteorol.* 239.

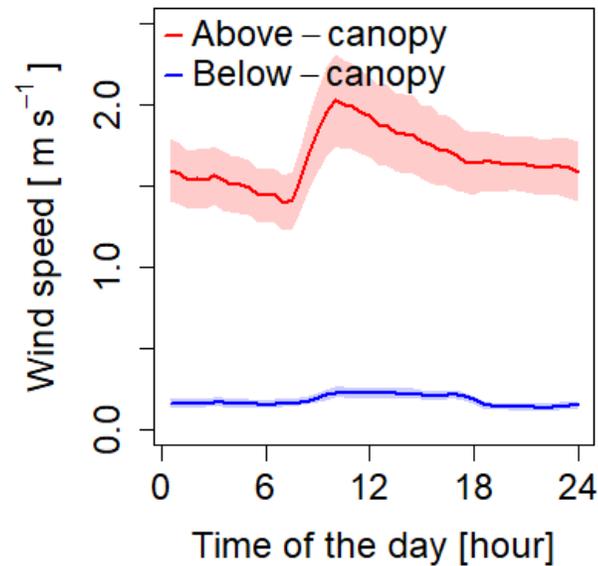
# Wind speed

- Wind speed is generally low in the oil palm plantation. Winds peak around noon and reach their minima shortly before sunrise
- Below the oil palm canopy, calm conditions strikingly dominate, especially during night

Daily average wind speed

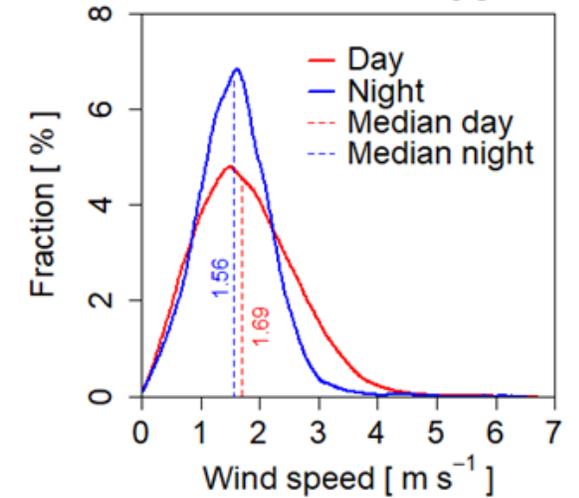


Diel horizontal wind speed

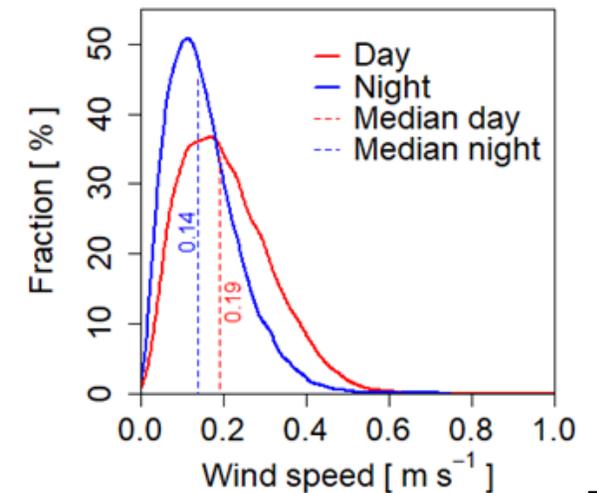


Probability density function of above- and below-canopy wind speed

Above-canopy



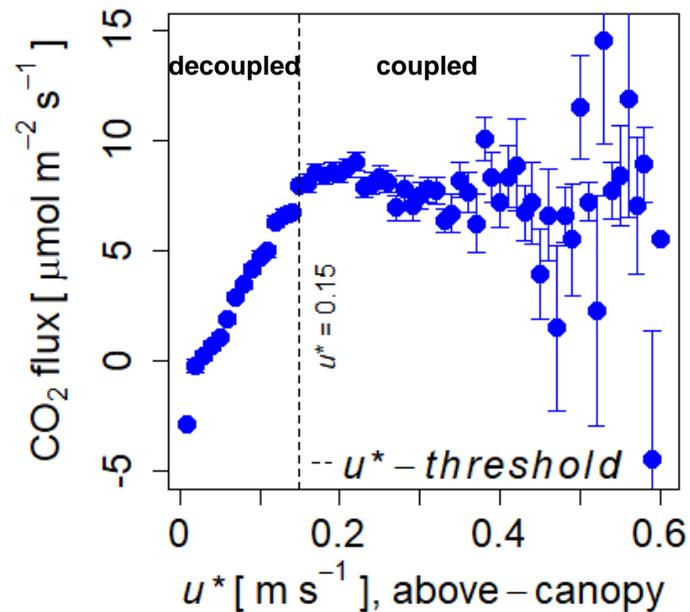
Below-canopy



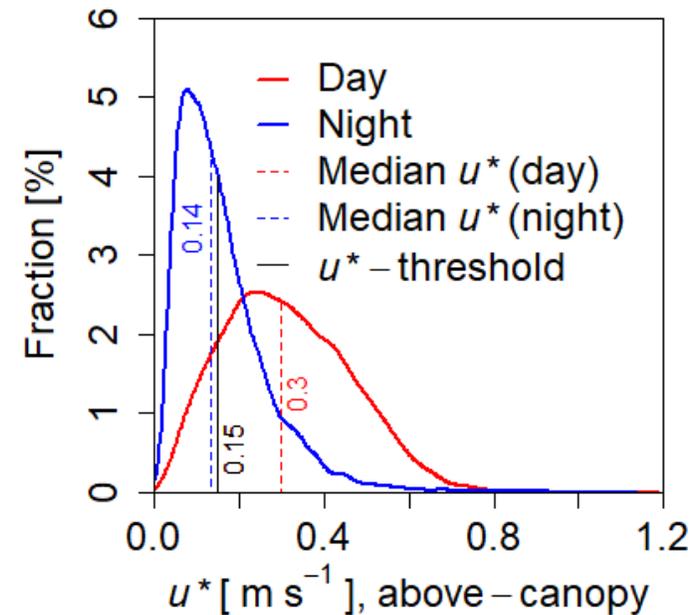
# Friction velocity ( $u^*$ )

- At higher nocturnal above-canopy turbulence strength ( $u^* > 0.15 \text{ m s}^{-1}$ ) nocturnal above-canopy  $\text{CO}_2$  flux levels off, indicating transition from decoupling to coupling at this threshold

Nocturnal above-canopy  
 $u^*$  and  $\text{CO}_2$  flux



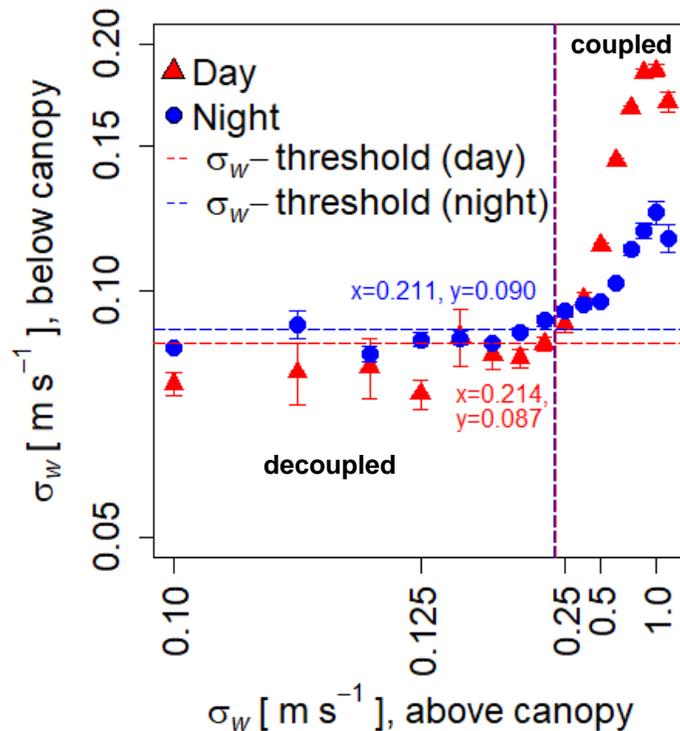
Probability density function  
of above-canopy  $u^*$



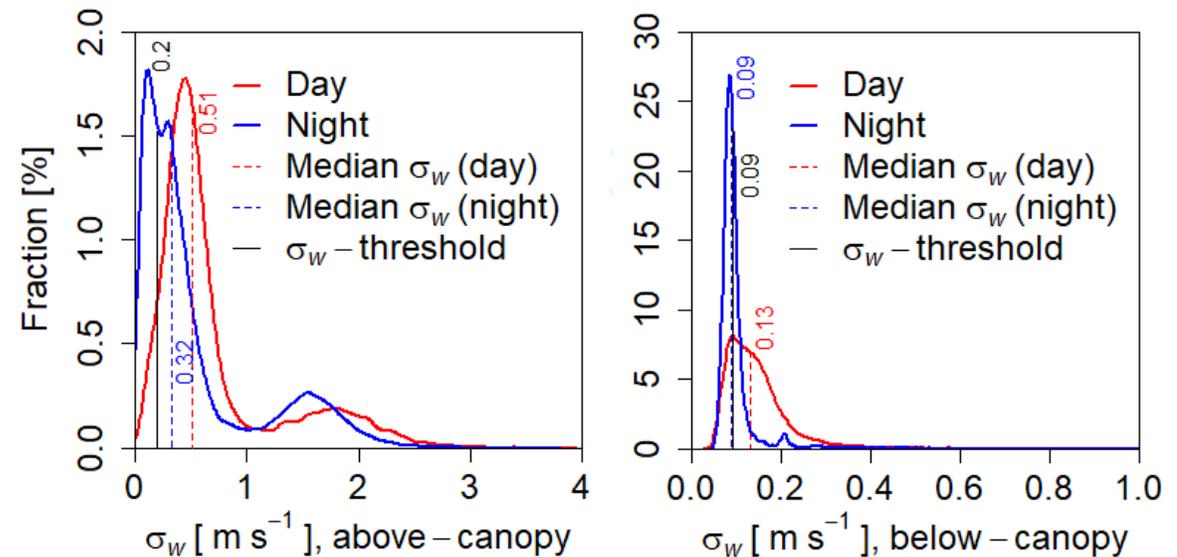
# Correlation between the standard deviation of below- vs. above-canopy vertical wind speed ( $\sigma_w$ )

- The correlation between above- and below-canopy  $\sigma_w$  breaks down at thresholds of  $\sim 0.21 \text{ m s}^{-1}$  and  $\sim 0.09 \text{ m s}^{-1}$ , respectively, indicating possible transition between decoupling to coupling between the two measurement heights at these thresholds

## Standard deviation of vertical wind speed

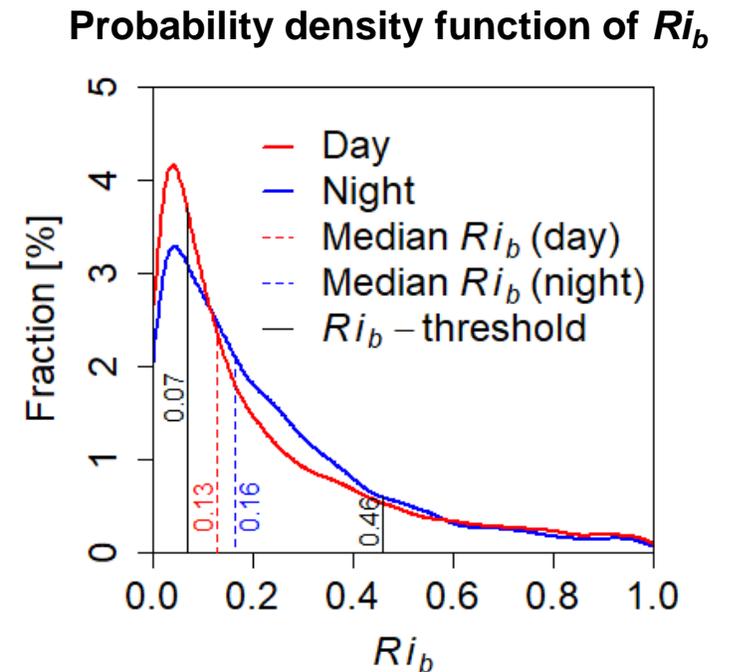
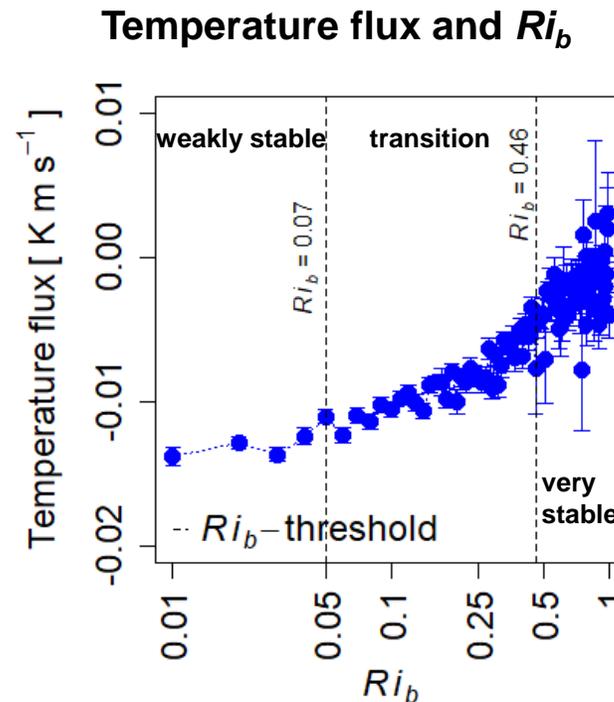
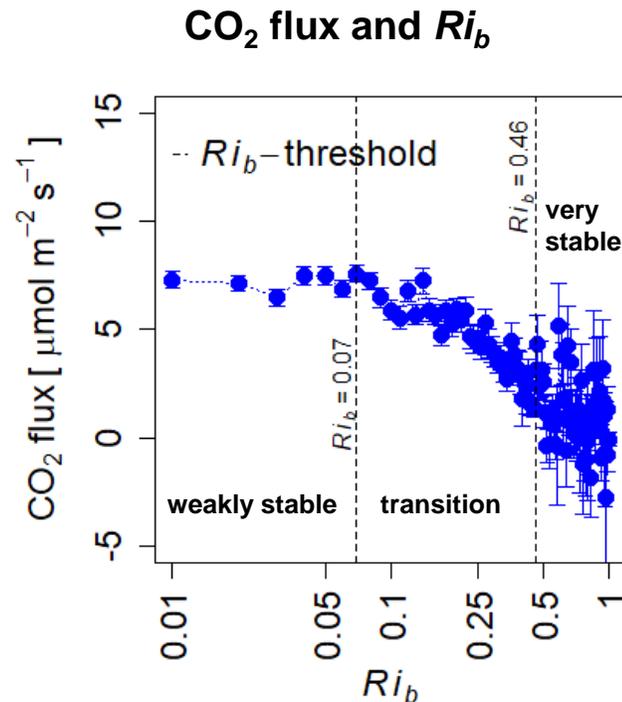


## Probability density function of above- and below-canopy $\sigma_w$



# Dynamic stability of the atmosphere (bulk Richardson number, $Ri_b$ )

- Based on the bulk Richardson number ( $Ri_b$ ) and its relationship between nocturnal  $\text{CO}_2$  net ecosystem exchange ( $NEE$ ) and temperature flux (kinematic heat flux) we found three distinct stable boundary-layer regimes in our data set
- Fluxes decrease rapidly with increasing  $Ri_b$  and non-turbulent flow becomes increasingly important, with decoupling occurring at very stable conditions



# Overview CO<sub>2</sub> flux filtering approaches (non-gapfilled)

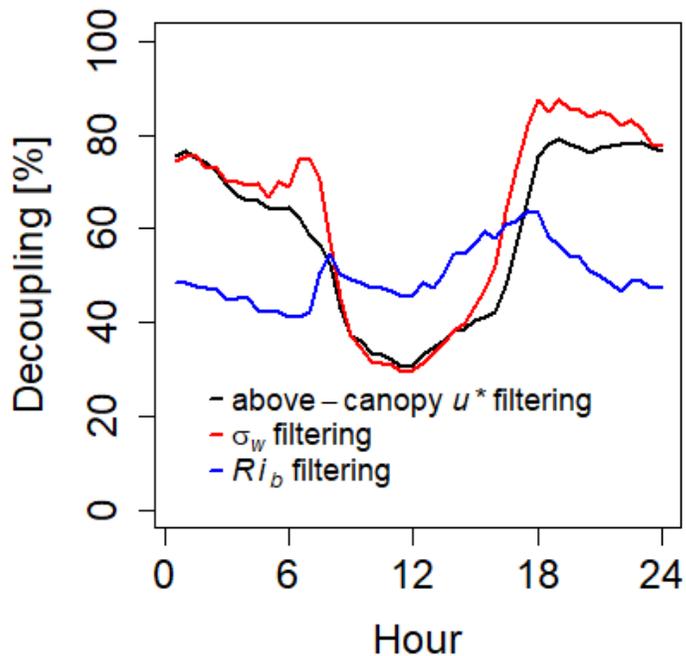
- On average, ~45% of CO<sub>2</sub> fluxes were removed due to filtering, most of them during night
- The  $\sigma_w$  filtering approach yielded highest daytime *NEE* (CO<sub>2</sub> uptake) while the  $u^*$ -filtering approach yielded highest nocturnal respiration

Filtering approach	Decoupling threshold	Percentage of CO <sub>2</sub> fluxes removed compared to unfiltered data (prior to gap filling)	CO <sub>2</sub> flux ± standard deviation [μmol m <sup>-2</sup> s <sup>-1</sup> ]			Coupling between above- and below-canopy air layer [%]	
			Average	Day	Night	Day	Night
Original data	-	-	-3.92 ± 14.66	-13.21 ± 13.15	6.05 ± 8.24	-	-
$u^*$ (friction velocity), above-canopy	$u^* < 0.15 \text{ m s}^{-1}$	33.5	-7.40 ± 15.73	-14.74 ± 12.61	8.12 ± 8.93	86.0	43.0
$\sigma_w$ (standard deviation of above- and below-canopy vertical wind speed)	<i>Above-canopy:</i> $\sigma_w < 0.214 \text{ m s}^{-1}$ (day) & $< 0.211 \text{ m s}^{-1}$ (night) <i>Below-canopy:</i> $\sigma_w < 0.087 \text{ m s}^{-1}$ (day) & $0.09 \text{ m s}^{-1}$ (night)	41.3	-9.19 ± 15.42	-16.43 ± 11.70	7.11 ± 9.05	93 (above-canopy), 80 (below-canopy)	68 (above-canopy), 44 (below-canopy)
$Ri_b$ (bulk Richardson number)	<i>Weakly stable:</i> $Ri_b < 0.07$ <i>Stable regime:</i> $Ri_b > 0.46$	59.8	-2.66 ± 14.63	-12.60 ± 13.50	6.52 ± 8.23	86.0	85.0

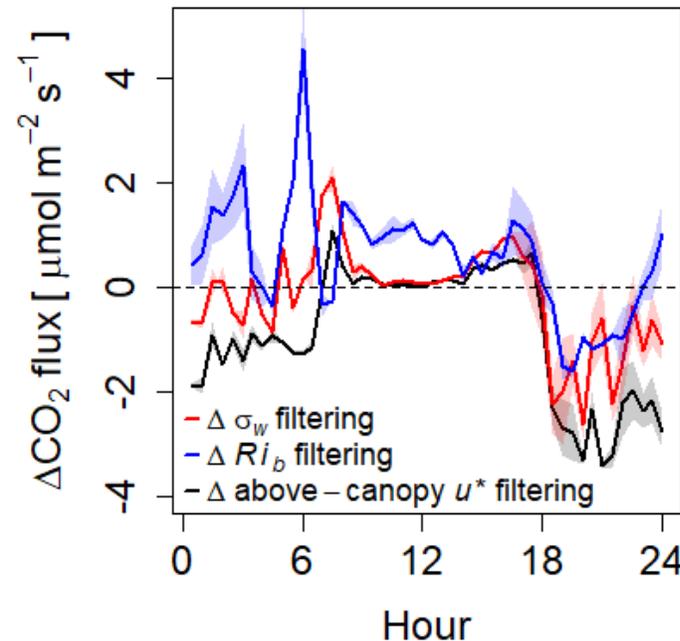
# Diel characteristics of decoupling and CO<sub>2</sub> fluxes, accumulated carbon

- The short periods after sunrise and before sunset are crucial for the breakdown and development of decoupling or coupling between above- and below-canopy air layers. They are also those times of the day (together with night conditions), when differences in *NEE* between the different flux filtering approaches are most apparent

Diel percentage of decoupling



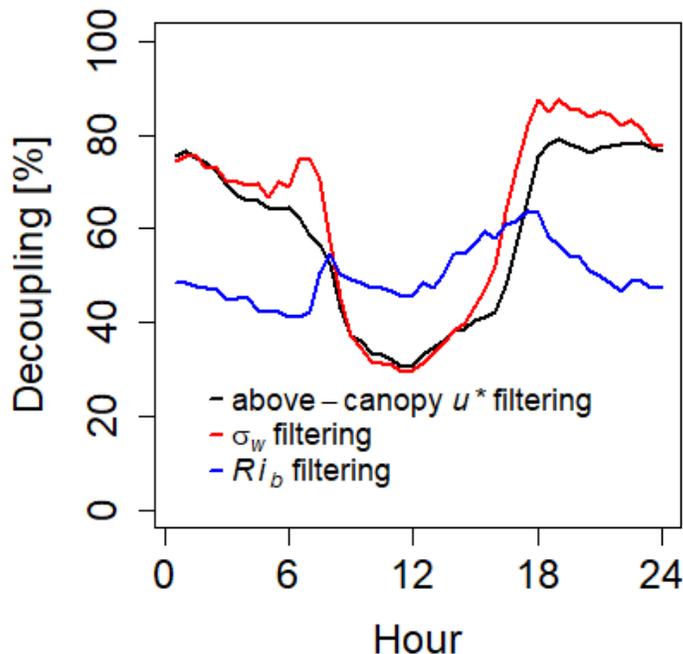
Absolute differences of CO<sub>2</sub> fluxes in relation to non-filtered data



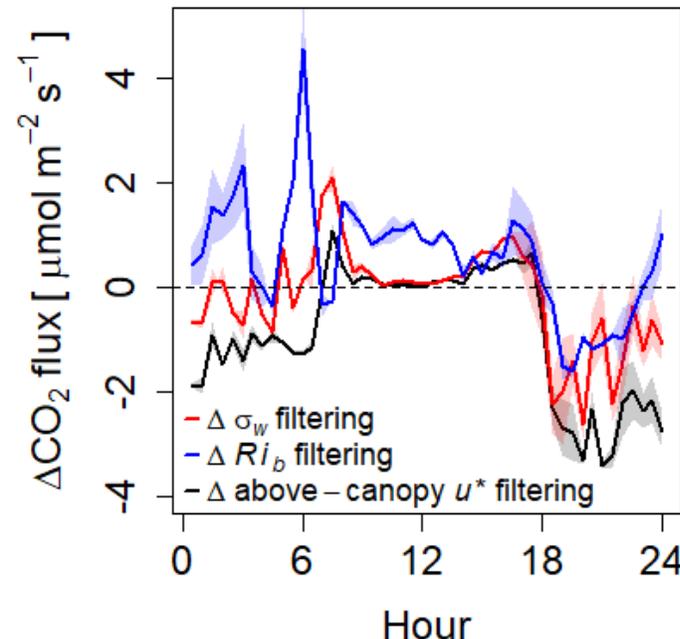
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- In 2019,  $\sigma_w$  filtered data yielded highest accumulated carbon while  $Ri_b$  filtered data yielded lowest accumulated carbon. Differences in accumulated carbon are up to 251.5 gC m<sup>-2</sup>

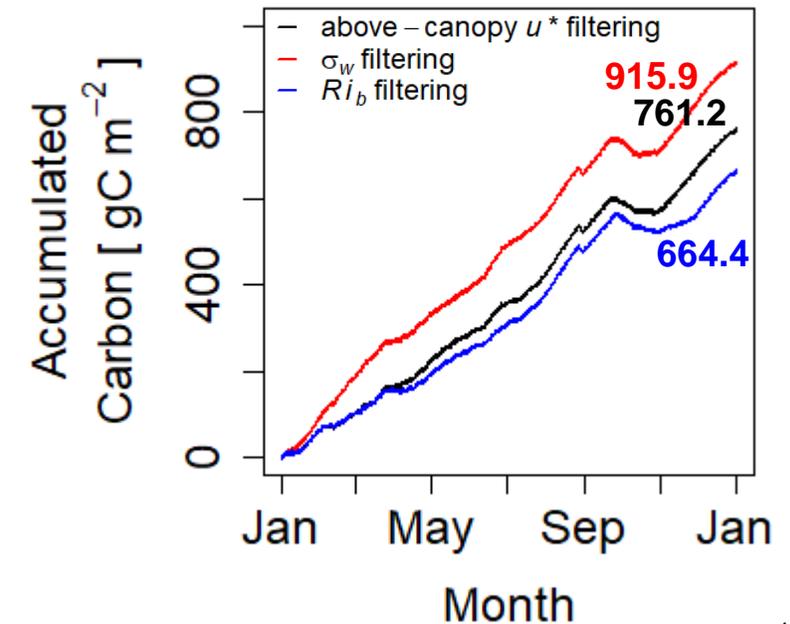
### Diel percentage of decoupling



### Absolute differences of CO<sub>2</sub> fluxes in relation to non-filtered data



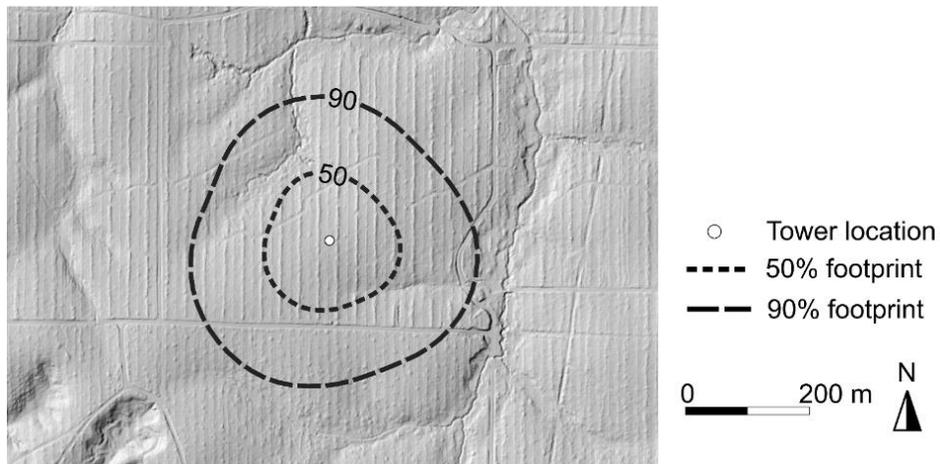
### Accumulated carbon (2019)



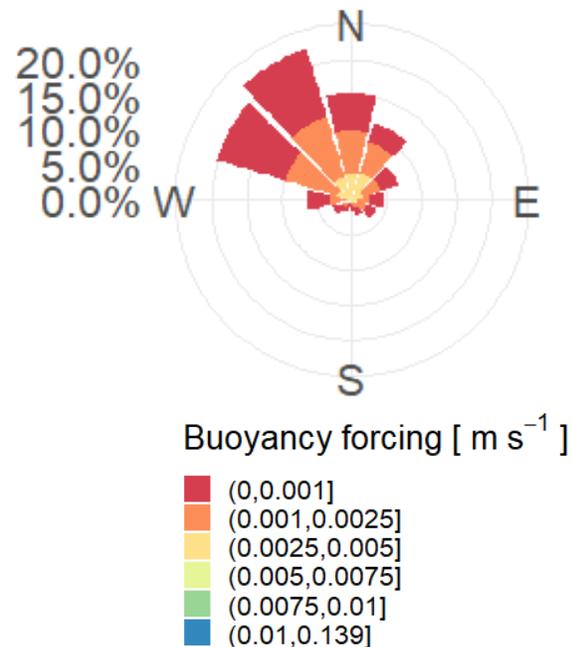
# Is the below-canopy flow topically induced?

- A preferential wind sector from northwest ( $\sim 315^\circ$ ) for below-canopy wind and buoyancy forcing is clearly visible
- Lowering terrain (from the tower's perspective in  $\sim 90^\circ$  to  $\sim 150^\circ$ ) and a slope of  $3^\circ$  within the footprint area of the *EC*-tower, may already be enough to create thermally-induced drainage flow

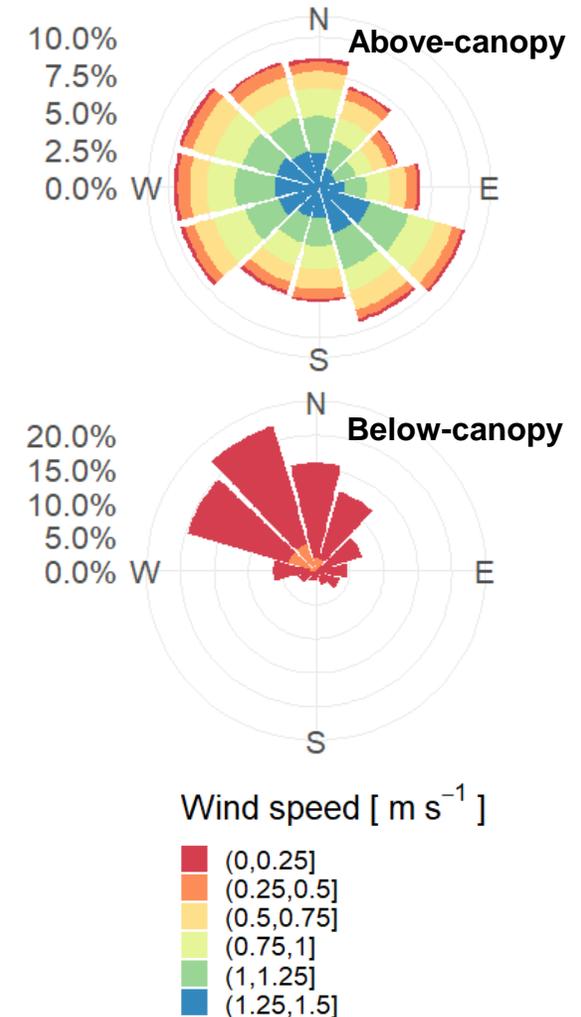
Hillshade map (from DEM) and EC footprint areas



Buoyancy forcing in dependency of below-canopy wind direction

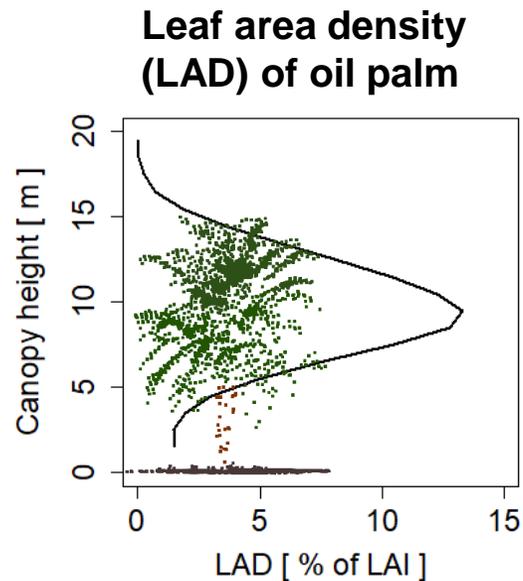


Wind roses for low above-canopy wind conditions and  $Ri_b > 0.07$

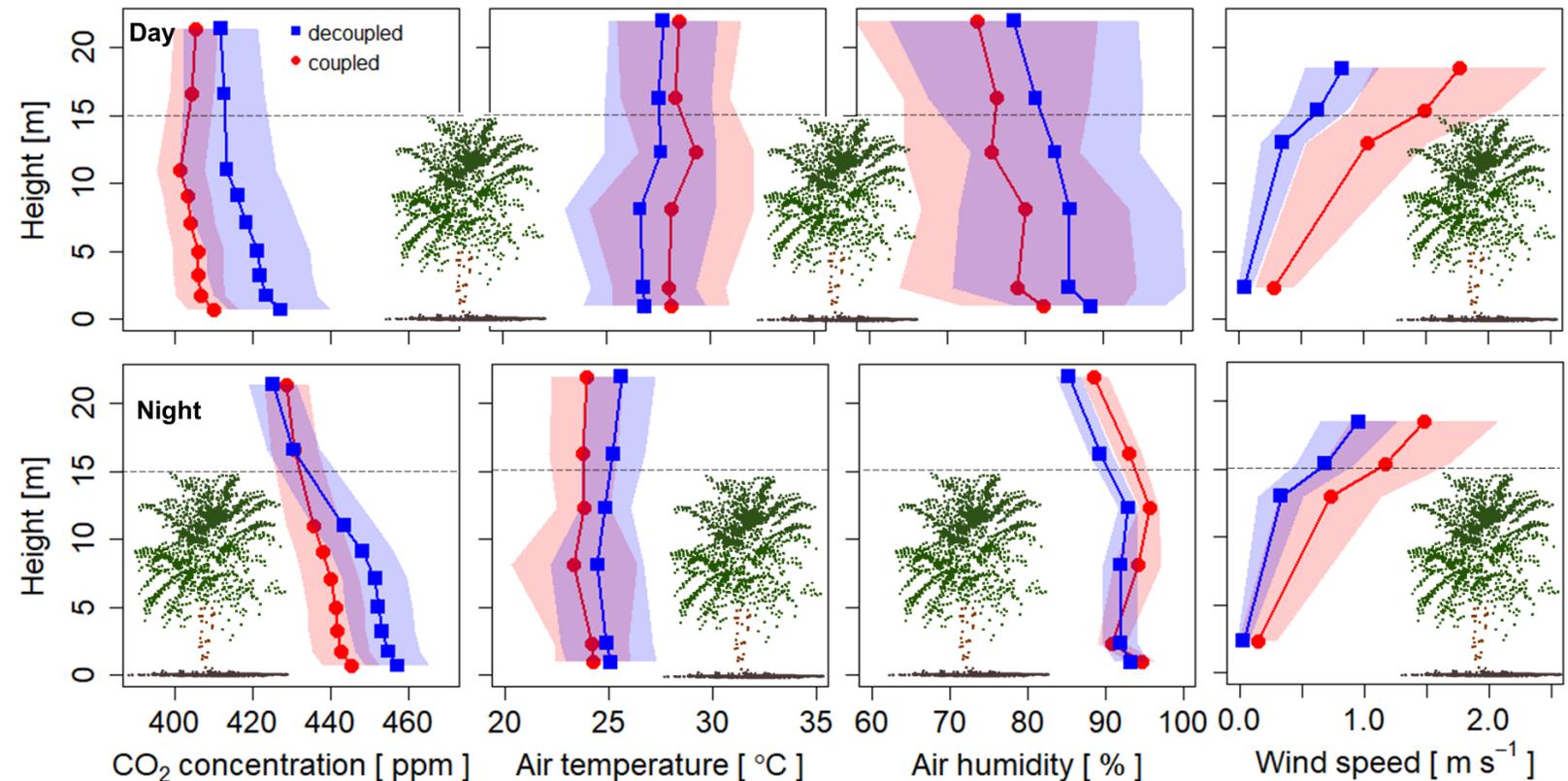


# Vertical structure of the air layer

- The oil palm canopy marks a clear boundary for the behavior of CO<sub>2</sub> concentration and wind
- The canopy is nearly always stably stratified, expressed by the temperature increase in the upper parts of the canopy. Under low-wind conditions, such stably stratified atmosphere potentially stimulates decoupling

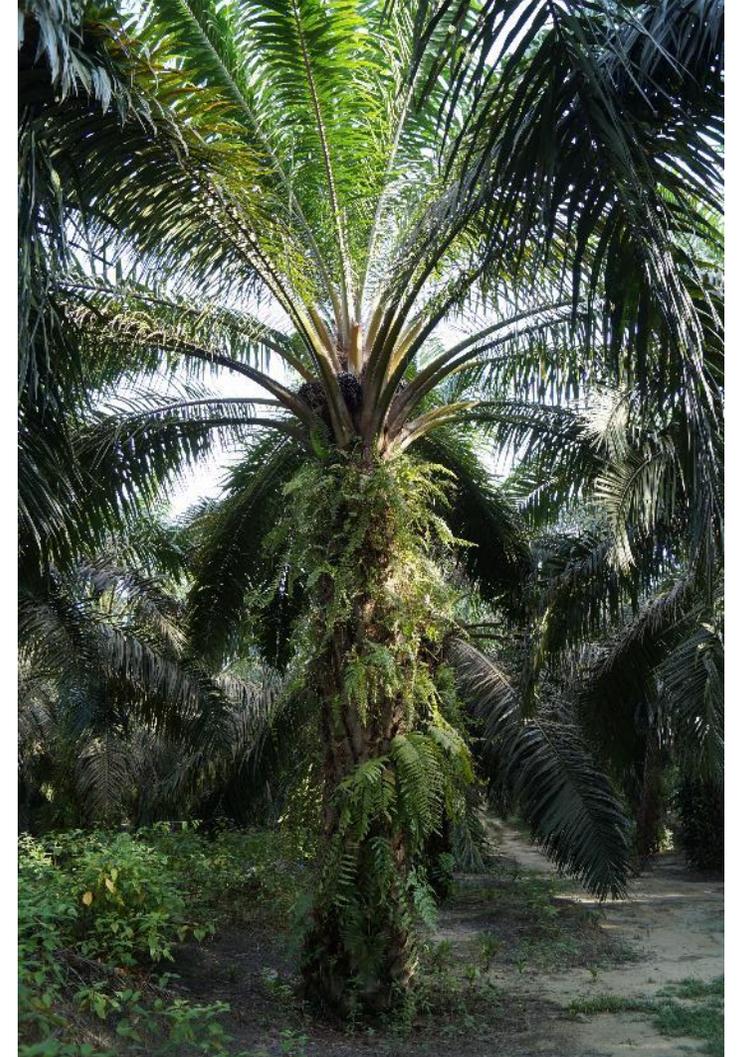


**Vertical profiles of CO<sub>2</sub> concentration and meteorological parameters**



# Summary & conclusion

- Wind speeds above and below the oil palm canopy are generally low
- Threshold analyses showed that during night, decoupling may occur frequently and up to 48% of measured nocturnal CO<sub>2</sub> fluxes fall within such decoupling periods
- Sensitivity in detection of decoupling is highly dependent on the applied method, with  $\sigma_w$  filter being most sensitive in the detection of decoupling
- CO<sub>2</sub> flux filtering approaches yield substantial differences in accumulated carbon
- A slope of 3° within the footprint area of the *EC*-tower, may already be enough to create thermally-induced drainage flow
- The canopy is nearly always stably stratified, expressed by the temperature increase in the upper parts of the canopy. Under low-wind conditions, such stably stratified atmosphere potentially may stimulate decoupling



# Take-home message

- Decoupling of above- and below-canopy air layers needs to be investigated and considered, especially in such low-wind tropical ecosystems and when eddy covariance data is used as reference for fluxes of tall vegetation or for modelling approaches

**THANK YOU FOR YOUR ATTENTION!**

**QUESTIONS?**



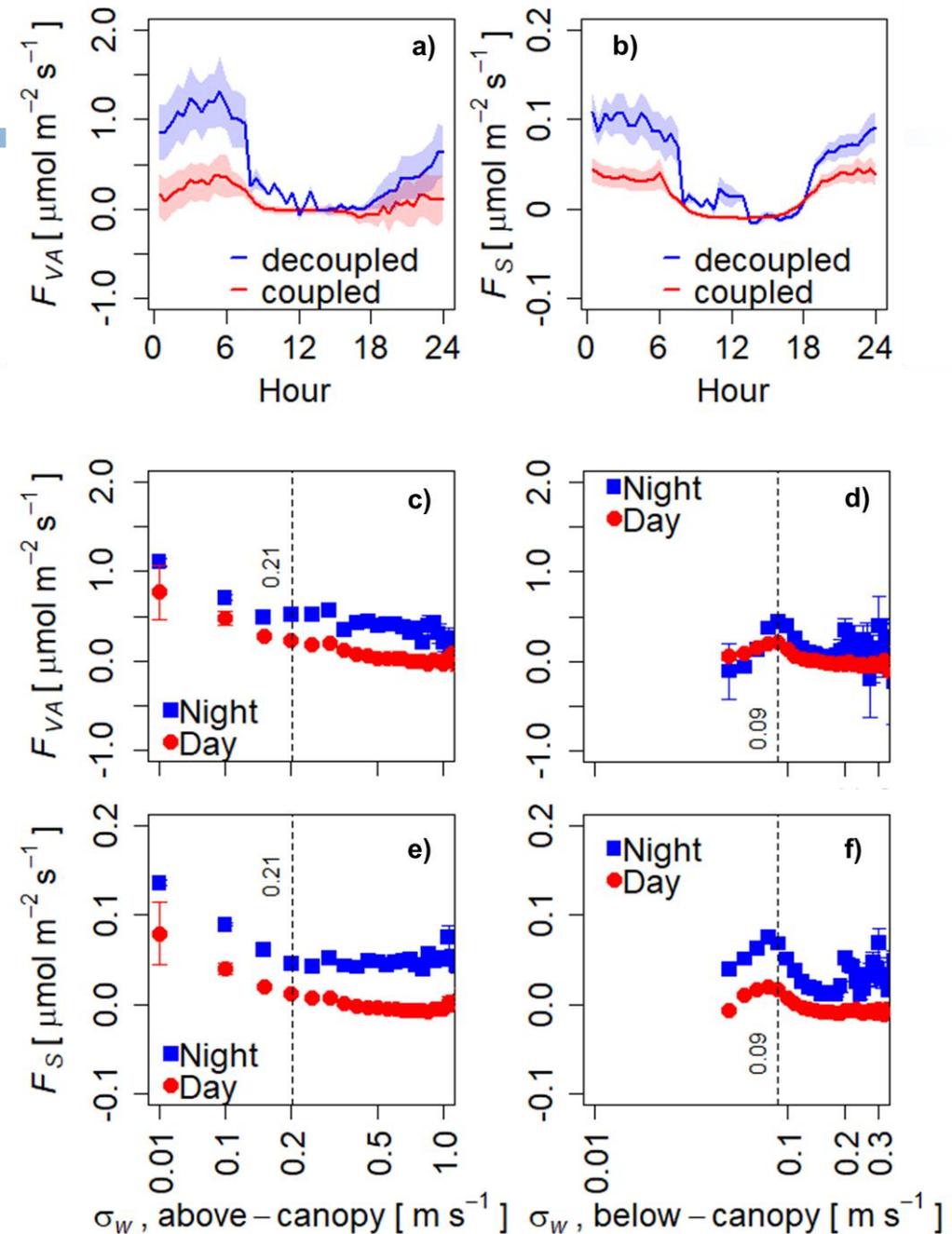
@BioclimGoe  
@efforts\_crc990  
@stiegler\_chr





# CO<sub>2</sub> vertical advection ( $F_{VA}$ ) and storage fluxes ( $F_S$ )

- The magnitude of  $F_{VA}$  and  $F_S$  peaks shortly before dawn, when CO<sub>2</sub> concentration reaches its diel maximum, and both vertical and horizontal wind speed approach their diel minima
- At low  $\sigma_w$ ,  $F_{VA}$  and  $F_S$  show a decreasing trend with increasing above-canopy  $\sigma_w$
- Below the canopy, an initial increase of  $F_{VA}$  and  $F_S$  at  $\sigma_w < \sigma_{w \text{ threshold}}$  may be linked to increasing, but still decoupled, below-canopy turbulent transport



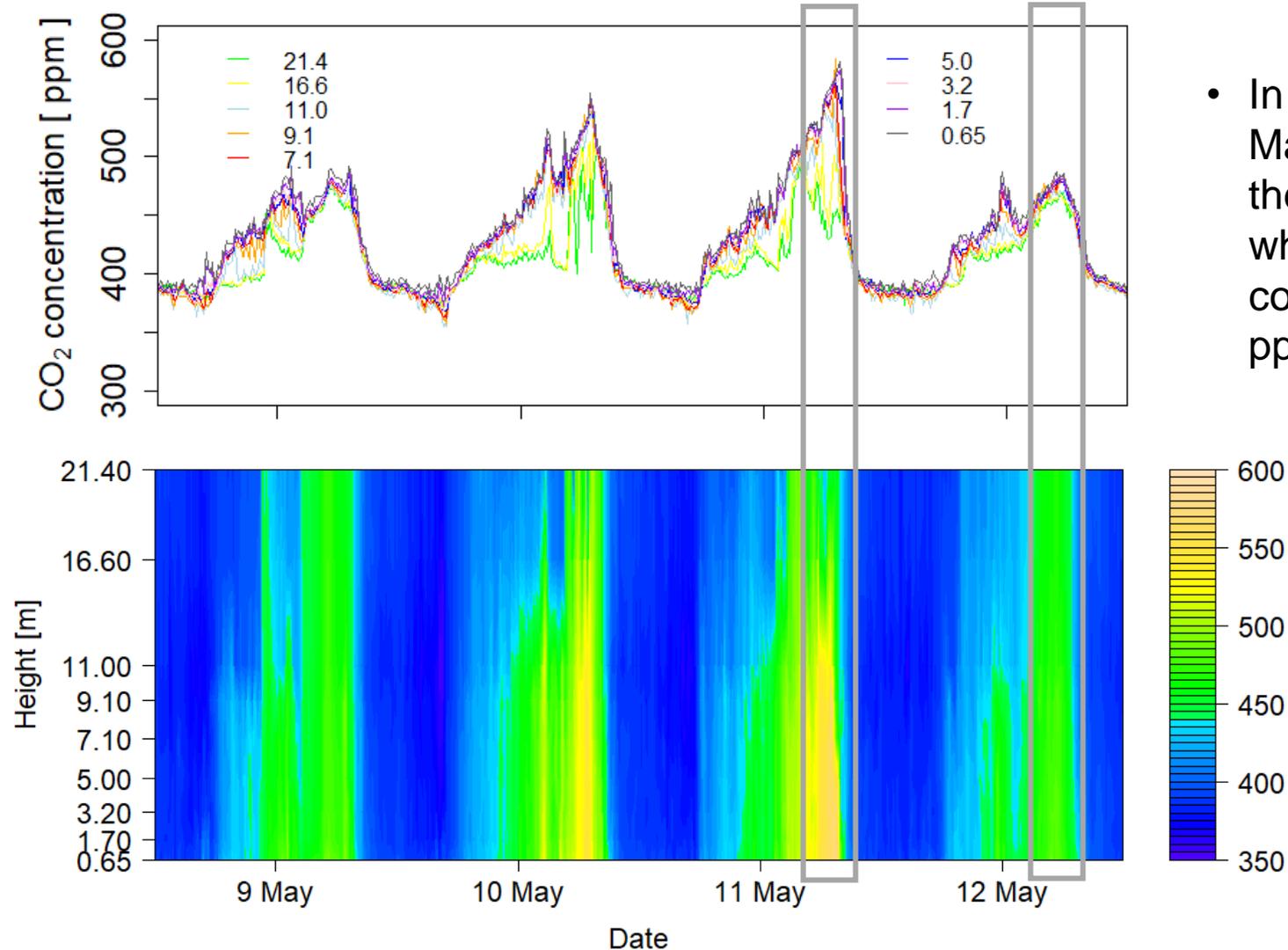
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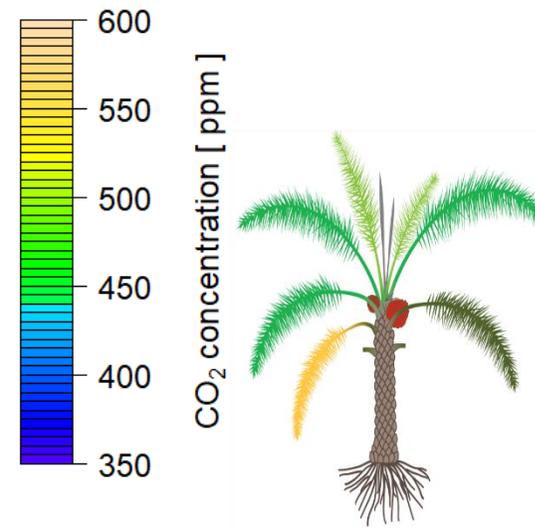
Filtering approach	Decoupling threshold	Percentage of CO <sub>2</sub> fluxes removed compared to unfiltered data (prior to gap filling)	CO <sub>2</sub> flux $\pm$ standard deviation [ $\mu\text{mol m}^{-2} \text{s}^{-1}$ ]			Coupling between above- and below-canopy air layer [%]	
			Average	Day	Night	Day	Night
<b><math>u^*</math></b> (friction velocity), above-canopy	$u^* < 0.15 \text{ m s}^{-1}$	33.5	$-2.48 \pm 14.17$	$-12.96 \pm 12.32$	$8.02 \pm 5.40$	86.0	43.0
<b><math>\sigma_w</math></b> (standard deviation of above- and below-canopy vertical wind speed)	<i>Above-canopy:</i> $\sigma_w < 0.214 \text{ m s}^{-1}$ (day) & $< 0.211 \text{ m s}^{-1}$ (night) <i>Below-canopy:</i> $\sigma_w < 0.087 \text{ m s}^{-1}$ (day) & $0.09 \text{ m s}^{-1}$ (night)	41.3	$-2.94 \pm 14.11$	$-13.64 \pm 11.89$	$7.77 \pm 5.27$	93 (above-canopy), 80 (below-canopy)	68 (above-canopy), 44 (below-canopy)
<b><math>Ri_b</math></b> (bulk Richardson number)	<i>Weakly stable:</i> $Ri_b < 0.07$ <i>Stable regime:</i> $Ri_b > 0.46$	59.8	$-2.19 \pm 13.91$	$-12.45 \pm 12.20$	$8.08 \pm 5.26$	86.0	85.0

# CO<sub>2</sub> concentration, 8 – 12 May 2019

CO<sub>2</sub> concentration, 8 May 2019 (12 h) to 12 May 2019, (12 h)



- In the morning hours of 10 and 11 May 2019, CO<sub>2</sub> concentrations below the canopy reach up to 580 ppm while above-canopy CO<sub>2</sub> concentration remains below 500 ppm.



CO<sub>2</sub> concentration profile during the period 8 May 2019 (12 h) to 12 May 2019, (12 h)

# CO<sub>2</sub> concentration & micrometeorological conditions, 8 – 12 May 2019

- Nights with high CO<sub>2</sub> concentration:
  - Calm conditions dominate in all vertical profiles
  - Air temperature is lower and air humidity is higher compared to nights with lower CO<sub>2</sub> concentration

CO<sub>2</sub> concentration, air temperature, air relative humidity, and wind speed profile during the period 8 May 2019 (12 h) to 12 May 2019, (12 h).

